

Modelling of shallow dispersive water waves

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“Numerical approximations of hyperbolic systems”
NumHyp2015
Cortona, Italy, June 14 – 20

Acknowledgements

to my precious collaborators

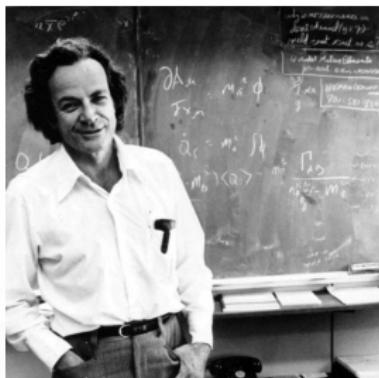
Collaborators:

- Didier CLAMOND: Professor (LJAD)
Université de Nice Sophia Antipolis, France
- Dimitrios MITSOTAKIS: Victoria University of
Wellington, New Zealand



Water wave problem - I

Prof. Richard FEYNMAN (1918 – 1988)



The Feynman Lectures on Physics (Vol. I)

[...] the next waves of interest, that are easily seen by everyone and which are usually used as an example of waves in elementary courses, are water waves. As we shall soon see, they are the worst possible example, because they are in no respects like sound and light; they have all the complications that waves can have [...]

Water wave problem - II

The mathematical formulation [1]

- Continuity equation
(incompressibility + irrotationality)

$$\nabla_{\mathbf{x},y}^2 \phi = 0, \quad (\mathbf{x}, y) \in \times [-d, \eta],$$

- Kinematic bottom condition:

$$\frac{\partial \phi}{\partial y} + \nabla \phi \cdot \nabla d = 0, \quad y = -d,$$

- Kinematic free surface condition:

$$\frac{\partial \eta}{\partial t} + \nabla \phi \cdot \nabla \eta = \frac{\partial \phi}{\partial y}, \quad y = \eta(\mathbf{x}, t),$$

- Dynamic free surface condition (Cauchy–Lagrange integral):

$$\frac{\partial \phi}{\partial t} + \frac{1}{2} |\nabla_{\mathbf{x},y} \phi|^2 + g\eta = 0, \quad y = \eta(\mathbf{x}, t).$$



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Dispersion relation analysis - I

Infinitesimal periodic waves on flat bottom

- Continuity equation
(incompressibility + irrotationality)

$$\nabla_{\mathbf{x},y}^2 \phi = 0, \quad (\mathbf{x}, y) \in \times [-d, 0],$$

- Kinematic bottom condition:

$$\frac{\partial \phi}{\partial y} = 0, \quad y = -d,$$

- Kinematic free surface condition:

$$\frac{\partial \eta}{\partial t} = \frac{\partial \phi}{\partial y}, \quad y = 0,$$

- Dynamic free surface condition (Cauchy–Lagrange integral):

$$\frac{\partial \phi}{\partial t} + g\eta = 0, \quad y = 0.$$



Dispersion relation analysis - II

Infinitesimal periodic waves on flat bottom

Plane wave solutions:

$$\eta(\mathbf{x}, t) = a e^{i(\mathbf{k} \cdot \mathbf{x} - \omega t)}, \quad \phi(\mathbf{x}, y, t) = b \varphi(y) e^{i(\mathbf{k} \cdot \mathbf{x} - \omega t)}.$$

Dispersion relation for gravity waves:

$$\omega^2(k) = gk \tanh(kd), \quad k := |\mathbf{k}| \in \mathbb{R}^+$$

Definition:

The wave propagation is:

- Dispersive: $c_p(k) = \frac{\omega(k)}{k} \neq \text{const}$
- Non-dispersive: $c_p(k) = \text{const}$ (e.g. sound, light, elastic waves)

Unfortunately, the water waves are dispersive, since

$$c_p(k) = \frac{\omega}{k} = \sqrt{gd \frac{\tanh(kd)}{kd}}$$

The case of Saint-Venant equations

Linearize and find plane wave solutions

This analysis is fundamentally linear!

- Linearized Saint-Venant equations:

$$\eta_t + d \nabla \cdot \mathbf{u} = 0,$$

$$\mathbf{u}_t + g \nabla \eta = 0.$$

- Plane wave solutions:

$$\eta(\mathbf{x}, t) = a e^{i(\mathbf{k} \cdot \mathbf{x} - \omega t)}, \quad \mathbf{u}(\mathbf{x}, t) = b e^{i(\mathbf{k} \cdot \mathbf{x} - \omega t)}$$



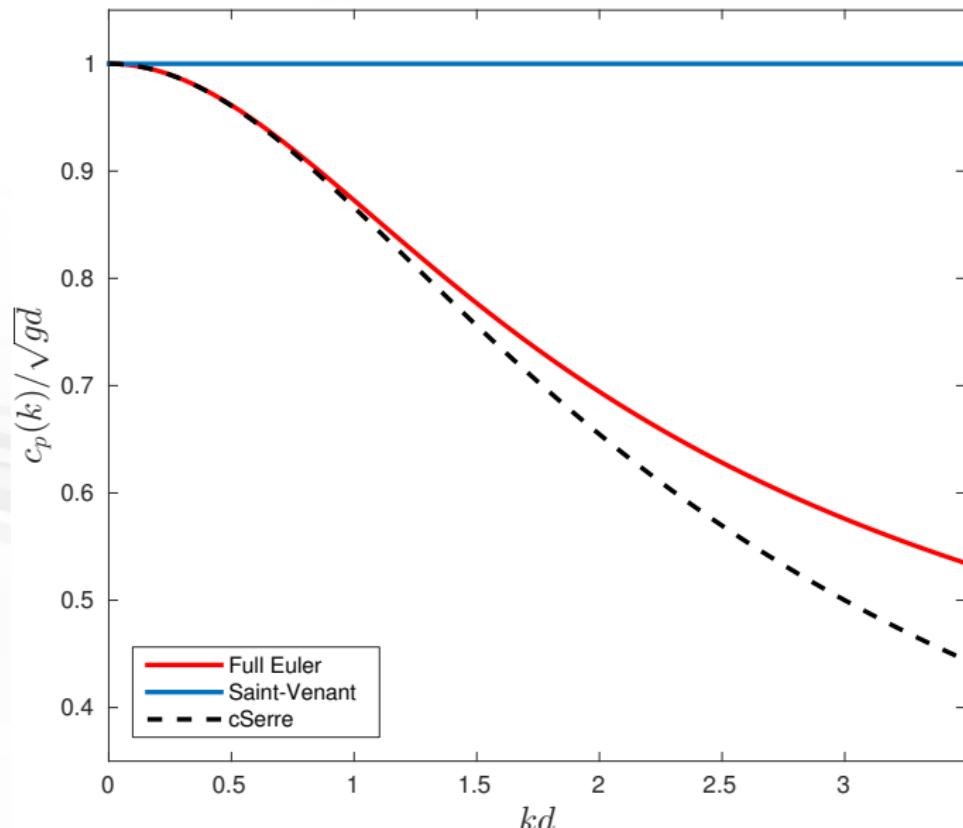
The dispersion relation is:

$$c_{\text{SV}} = \frac{\omega}{k} = \sqrt{gd} \equiv \lim_{kd \rightarrow 0} c_{\text{Euler}}(k)$$

The Saint-Venant equations are dispersionless!

Comparison of dispersion relations - I

for models considered so far...



Experimental evidences

The classical dam-break problem

- Courtesy of V. I. BUKREEV *et al.* [2]:

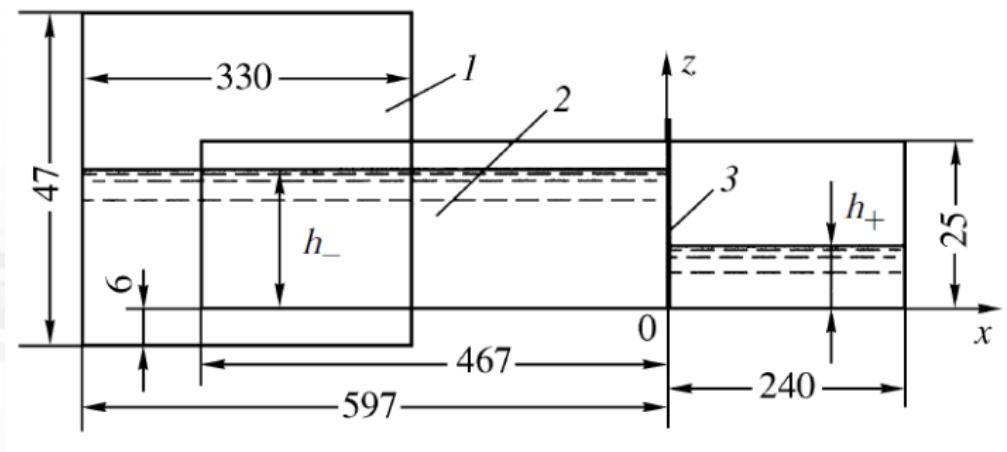


Figure: Experimental set-up.

Experimental evidences

The classical dam-break problem

- Courtesy of V. I. BUKREEV *et al.* [2]:

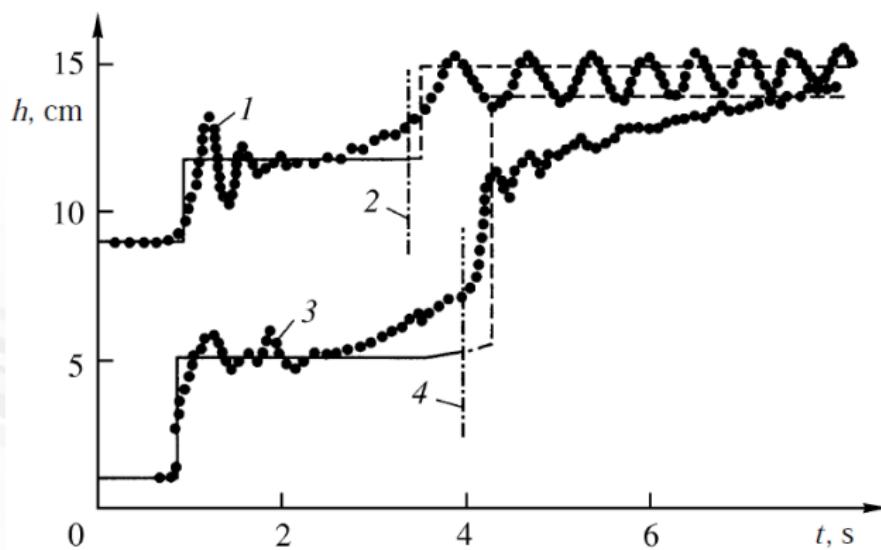


Figure: Wave height records.

Serre–(Green–Naghdi) equations

Shallow water equations

Main constitutive assumptions:

- Free surface is a graph: $y = \eta(\mathbf{x}, t)$
- Long wave: $\lambda \gg d$
- Nonlinearity is finite: $\varepsilon \equiv a/d \propto \mathcal{O}(1)$



Credits:

- Lord RAYLEIGH (1876) [3] (only *steady* version)
- François SERRE (1953) [4]
- C. SU & C. GARDNER (1969) [5]
- A. GREEN & P. NAGHDI (1976) [6]
- E. PELINOVSKY & ZHELEZNYAK (1985) [7]
- ...
- Nora AISSIOUENE. (15 June 2015) ☺



Paul M. Naghdi

Derivation of Rayleigh–Serre–Green–Naghdi equations - I

Depth-averaged shallow water ansatz

- Total water depth: $h(x, t) := d + \eta(x, t)$

Assumption:

- Depth-averaged profile:

$$u(x, y, t) \approx \bar{u}(x, t) \equiv \frac{1}{h} \int_{-d}^{\eta} u(x, y, t) dy$$

Incompressibility yields:

$$u_x + v_y = 0 \implies v(x, y, t) \approx -(y + d)\bar{u}_x$$

The energy can be computed:

Kinetic: $\frac{\mathcal{K}}{\rho} = \frac{1}{2} \int_{-d}^{\eta} (u^2 + v^2) dy \approx \frac{1}{2} h \bar{u}^2 + \frac{1}{6} h^3 \bar{u}_x^2$

Potential: $\frac{\mathcal{V}}{\rho} = \int_{-d}^{\eta} g(y + d) dy = \frac{1}{2} g h^2$

- The action is:

$$\frac{\mathcal{S}}{\rho} = \int_{t_1}^{t_2} \int_{x_1}^{x_2} \left[\frac{h \bar{u}^2}{2} + \frac{h^3 \bar{u}_x^2}{6} - \frac{g h^2}{2} + \{ h_t + [h \bar{u}]_x \} \tilde{\phi} \right] dx dt$$

- Euler–Lagrange equations:

$$\delta \tilde{\phi} : 0 = h_t + [h \bar{u}]_x,$$

$$\delta \bar{u} : 0 = \tilde{\phi} h_x - [h \tilde{\phi}]_x - \frac{1}{3} [h^3 \bar{u}_x]_x + h \bar{u},$$

$$\delta h : 0 = \frac{1}{2} \bar{u}^2 - g h + \frac{1}{2} h^2 \bar{u}_x^2 - \tilde{\phi}_t + \tilde{\phi} \bar{u}_x - [\bar{u} \tilde{\phi}]_x.$$

- After some simple algebra:

$$\tilde{\phi}_x = \bar{u} - \frac{1}{3} h^{-1} [h^3 \bar{u}_x]_x,$$

$$\tilde{\phi}_t = \frac{1}{2} h^2 \bar{u}_x^2 - \frac{1}{2} \bar{u}^2 - g h + \frac{1}{3} \bar{u} h^{-1} [h^3 \bar{u}_x]_x.$$

$$\implies \tilde{\phi}_{tx} = \tilde{\phi}_{xt} !$$

Derivation of Rayleigh–Serre–Green–Naghdi equations - III

Together with its conservation laws

- Mass conservation:

$$h_t + [h\bar{u}]_x = 0$$

- Momentum conservation:

$$[h\bar{u}]_t + \left[h\bar{u}^2 + \frac{1}{2}gh^2 + \frac{1}{3}h^2\tilde{\gamma} \right]_x = 0$$

$$\tilde{\gamma} \equiv h \left[\bar{u}_x^2 - \bar{u}_{xt} - \bar{u}\bar{u}_{xx} \right] = 2h\bar{u}_x^2 - h\partial_x[\bar{u}_t + \bar{u}\bar{u}_x]$$

- Surface tangential momentum:

$$\left[\bar{u} - \frac{1}{3}h^{-1}(h^3\bar{u}_x)_x \right]_t + \left[\frac{1}{2}\bar{u}^2 + gh - \frac{1}{2}h^2\bar{u}_x^2 - \frac{1}{3}\bar{u}h^{-1}(h^3\bar{u}_x)_x \right]_x = 0$$

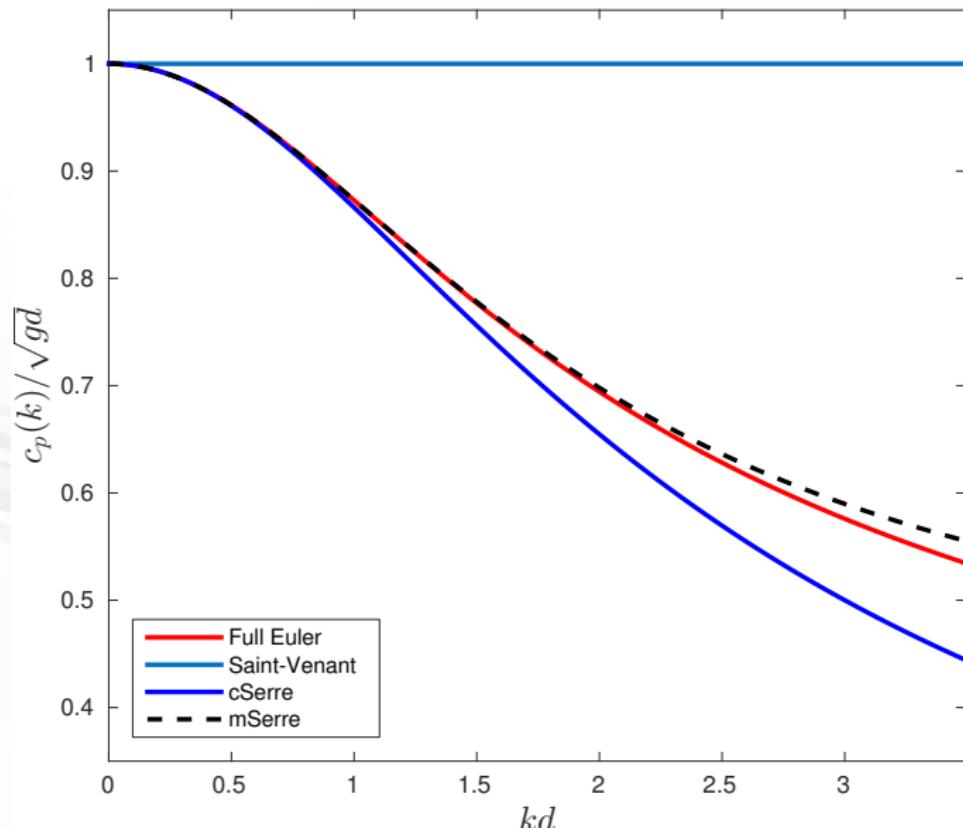
- Energy conservation:

$$\left[\frac{1}{2}h\bar{u}^2 + \frac{1}{6}h^3\bar{u}_x^2 + \frac{1}{2}gh^2 \right]_t +$$

$$\left[\left(\frac{1}{2}\bar{u}^2 + \frac{1}{6}h^2\bar{u}_x^2 + gh + \frac{1}{3}h\gamma \right) h\bar{u} \right]_x = 0$$

Comparison of dispersion relations - II

look at the black dotted line



Modified Serre's equations - I

Introduction of a free parameter into the model

- Introduction of scaled variables ($\partial_{x,t} = \mathcal{O}(\sigma)$):

$$x^* = \sigma x, \quad t^* = \sigma t, \quad \tilde{\gamma}^* = \sigma^{-2} \tilde{\gamma}$$

- Derived above vertical acceleration at free surface:

$$\tilde{\gamma}_{\text{old}} = 2\sigma^2 h \bar{u}_{x^*}^2 - \sigma^2 h \partial_{x^*} [\bar{u}_{t^*} + \bar{u} \bar{u}_{x^*}]$$

- Non-conservative form of momentum equation:

$$\bar{u}_{t^*} + \bar{u} \bar{u}_{x^*} = -g h_{x^*} - \sigma^2 \frac{1}{3} h^{-1} \partial_{x^*} [h^2 \tilde{\gamma}^*],$$

- New asymptotically equivalent acceleration:

$$\tilde{\gamma}_{\text{new}} = \sigma^2 2 h \bar{u}_{x^*}^2 + \sigma^2 g h h_{x^* x^*} + \mathcal{O}(\sigma^4),$$

- Bona-Smith [8] & Nwogu [9] trick ($0 \leq \alpha \leq 1$):

$$\begin{aligned} \tilde{\gamma}_\alpha := \alpha \tilde{\gamma}_{\text{old}} + (1 - \alpha) \tilde{\gamma}_{\text{new}} &= 2\sigma^2 h \bar{u}_{x^*}^2 + \\ (1 - \alpha) \sigma^2 g h h_{x^* x^*} - \alpha \sigma^2 h \partial_{x^*} [\bar{u}_{t^*} + \bar{u} \bar{u}_{x^*}] &+ \mathcal{O}(\sigma^4) \end{aligned}$$

Modified Serre's equations - II

The system of governing equations

- Modified Serre–(Green–Naghdi) equations:

$$h_t + [h\bar{u}]_x = 0,$$

$$[h\bar{u}]_t + \left[h\bar{u}^2 + \frac{1}{2}gh^2 + \frac{1}{3}h^2\tilde{\gamma}_\alpha \right]_x = 0.$$

$$\tilde{\gamma}_\alpha = 2h\bar{u}_x^2 + (1-\alpha)ghh_{xx} - \alpha h\partial_x[\bar{u}_t + \bar{u}\bar{u}_x]$$

Dispersion relation analysis:

- mSerre system:

$$\frac{c^2}{gd} = \frac{3 + (\alpha - 1)(kd)^2}{3 + \alpha(kd)^2} = 1 - \frac{1}{3}(kd)^2 + \frac{1}{9}\alpha(kd)^4 - \frac{1}{27}\alpha^2(kd)^6 + \dots$$

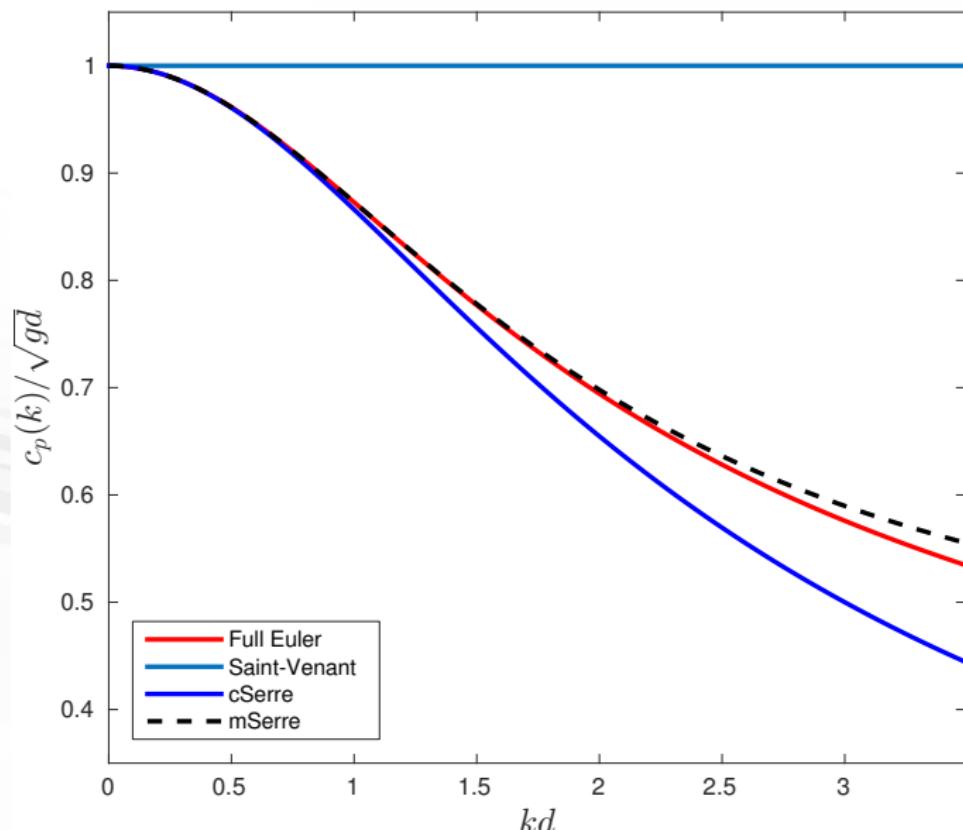
- Full Euler:

$$\frac{c^2}{gd} = \frac{\tanh(kd)}{kd} = 1 - \frac{1}{3}(kd)^2 + \frac{2}{15}(kd)^4 - \frac{17}{315}(kd)^6 + \dots$$

An *optimal* value: $\alpha = 6/5 = 1.2 !$

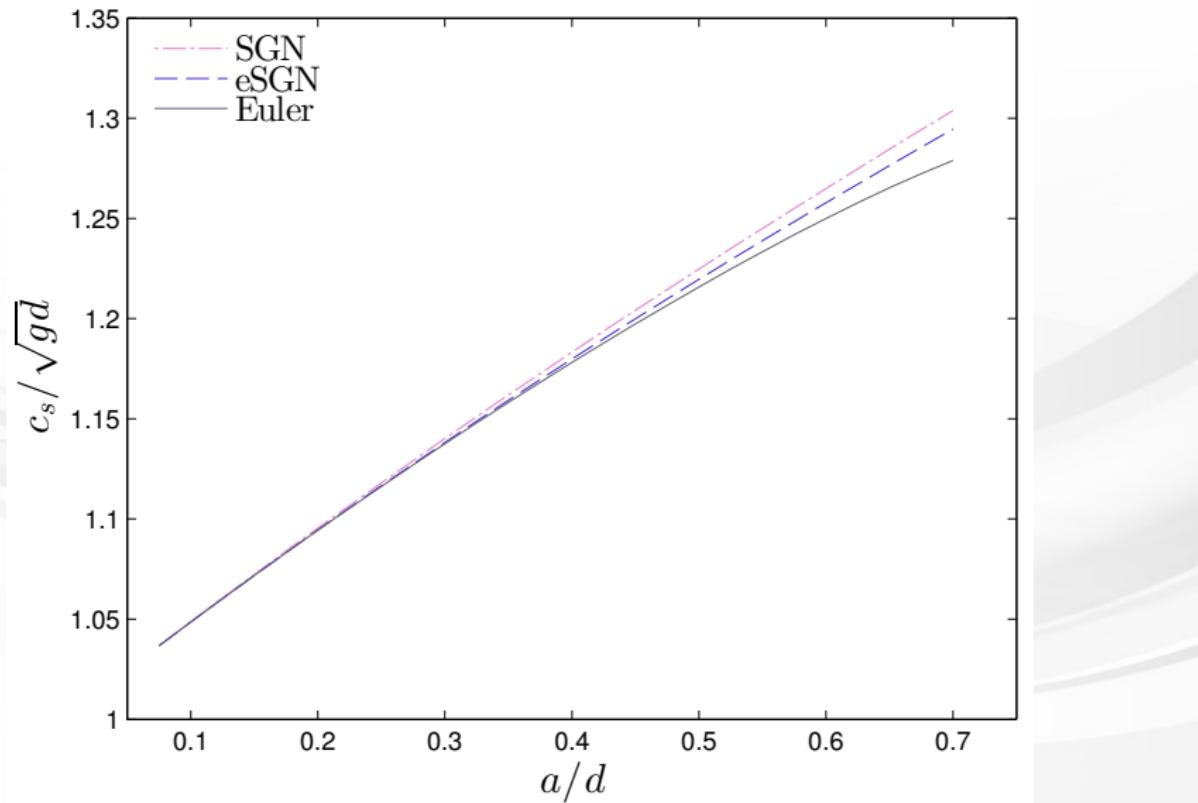
Comparison of dispersion relations - II

look at the black dotted line



Performance of the modified Serre model - I

Speed-amplitude relation [10]



Performance of the modified Serre model - II

Comparison of the classical, modified Serre and the full Euler equations [10]

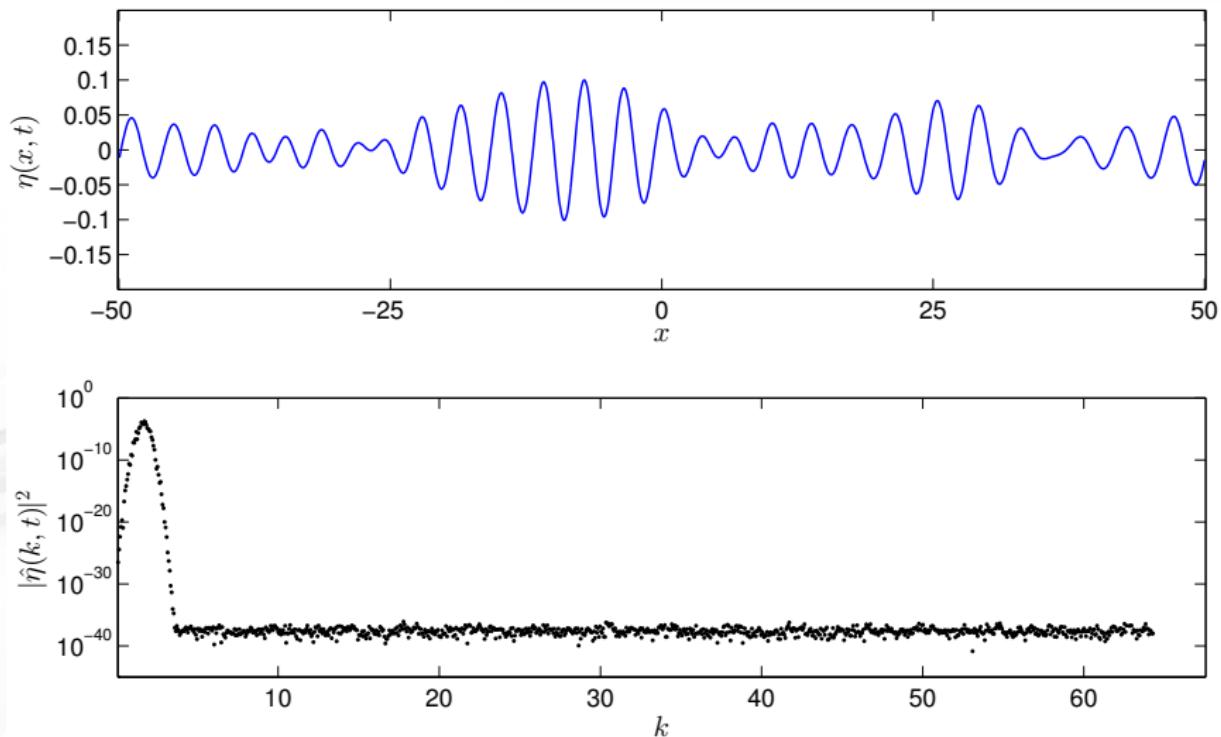


Figure: Initial condition (up) along with its Fourier power spectrum (down).

Performance of the modified Serre model - II

Comparison of the classical, modified Serre and the full Euler equations [10]

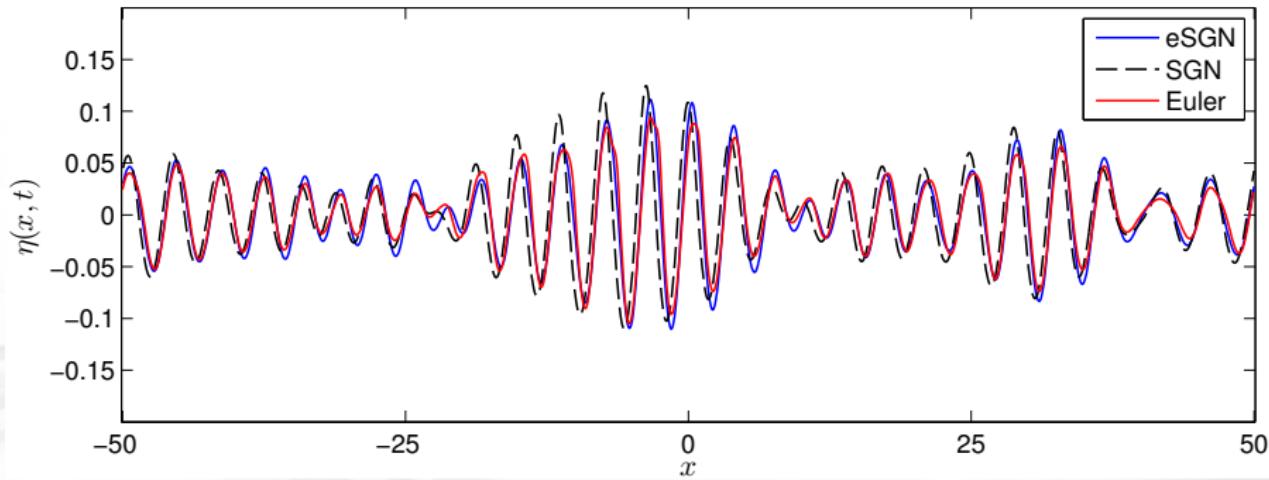


Figure: Free surface elevations at $t = 10\sqrt{d/g}$.

Performance of the modified Serre model - II

Comparison of the classical, modified Serre and the full Euler equations [10]

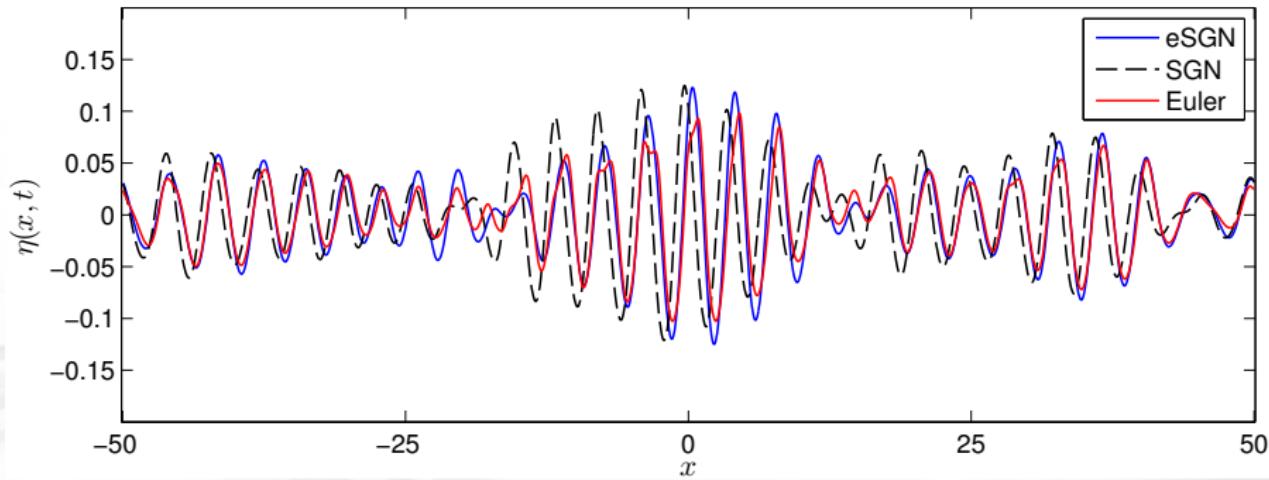


Figure: Free surface elevations at $t = 20\sqrt{d/g}$.

Performance of the modified Serre model - II

Comparison of the classical, modified Serre and the full Euler equations [10]

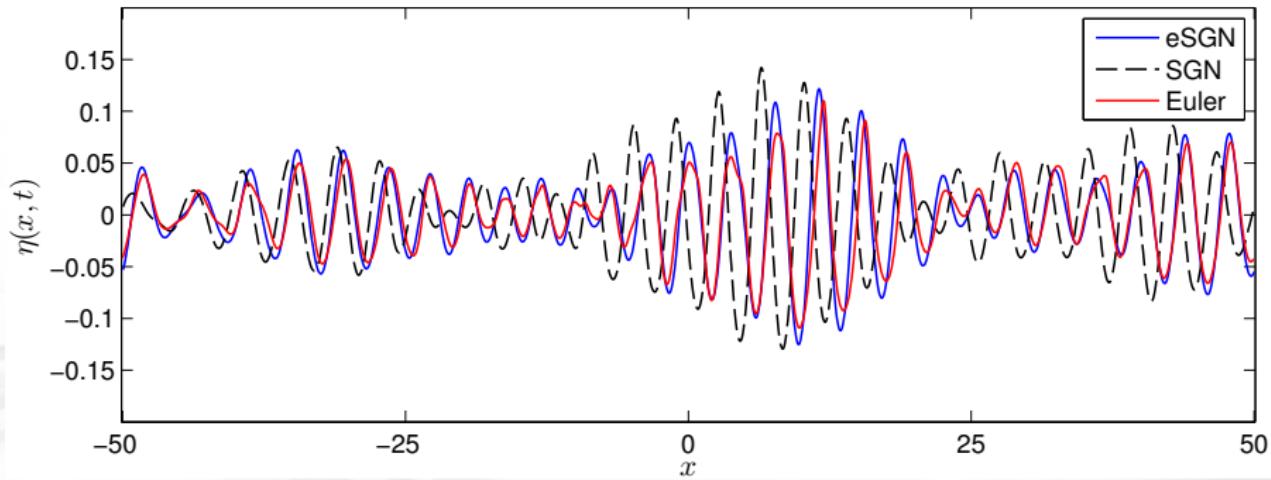


Figure: Free surface elevations at $t = 40\sqrt{d/g}$.

Performance of the modified Serre model - II

Comparison of the classical, modified Serre and the full Euler equations [10]

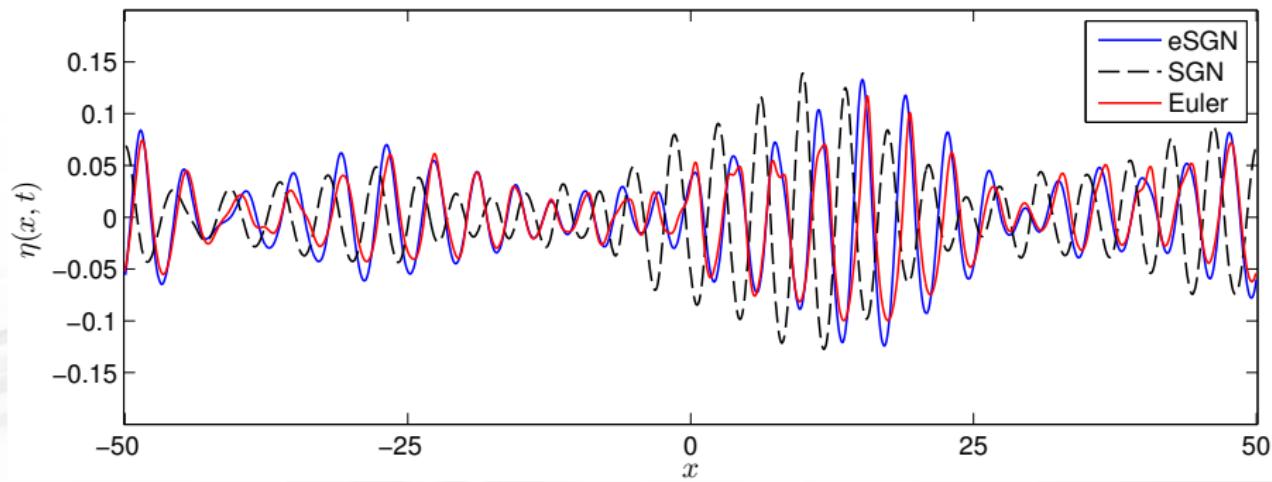


Figure: Free surface elevations at $t = 50\sqrt{d/g}$.

Performance of the modified Serre model - II

Comparison of the classical, modified Serre and the full Euler equations [10]

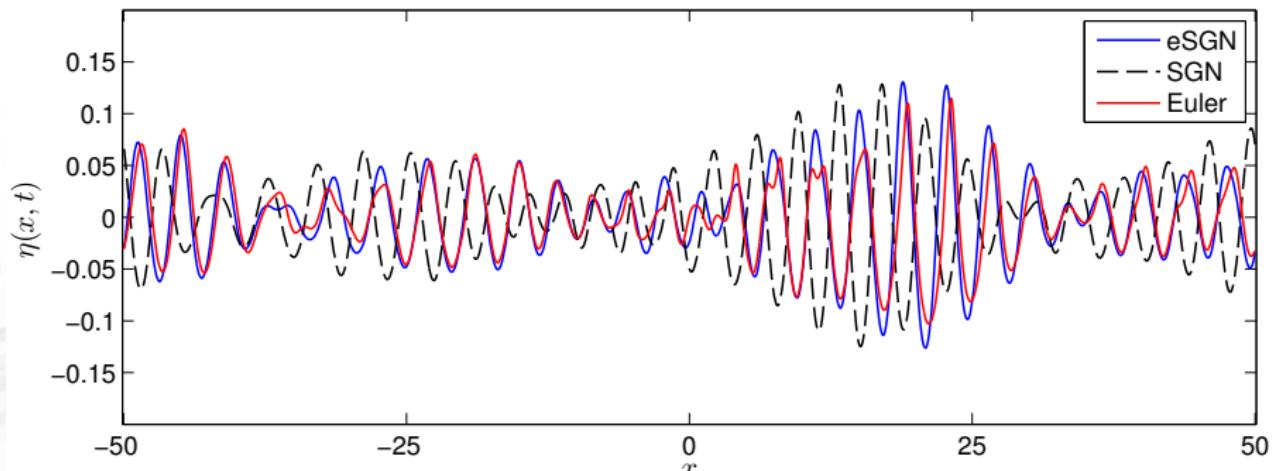


Figure: Free surface elevations at $t = 60\sqrt{d/g}$.

Performance of the modified Serre model - II

Comparison of the classical, modified Serre and the full Euler equations [10]

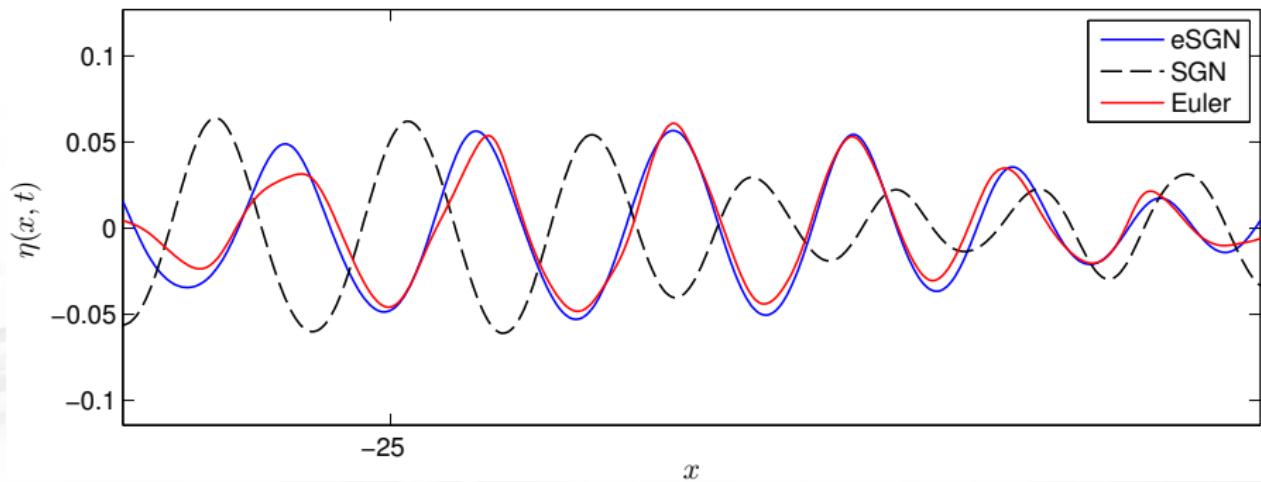


Figure: Zoom on free surface elevations at $t = 60\sqrt{d/g}$.

The mSerre system is not perfect (as well as Nwogu's system)!

There is at least one important issue

- Galilean invariance: **OK** ☺
- Energy conservation: presumably **NO** ☹



Derivation of a consistent mSerre model - I

Idea: manipulate the Lagrangian density!

- The Lagrangian density is:

$$\mathcal{L} = \frac{h \bar{u}^2}{2} + \frac{h^3 \bar{u}_x^2}{6} - \frac{g h^2}{2} + \{ h_t + [h \bar{u}]_x \} \tilde{\phi}$$

- Introduce explicitly the vertical acceleration:

$$\mathcal{L} = \frac{h \bar{u}^2}{2} + \frac{h^2 \tilde{\gamma}}{12} + \frac{h^3}{12} [\bar{u}_t + \bar{u} \bar{u}_x]_x - \frac{g h^2}{2} + \{ h_t + [h \bar{u}]_x \} \tilde{\phi}$$

- Substitute the relaxed version $\tilde{\gamma}_\alpha$ and simplify ($\beta \equiv 1 - \alpha$):

$$\mathcal{L}' = \frac{h \bar{u}^2}{2} + \frac{(2 + 3\beta) h^3 \bar{u}_x^2}{12} - \frac{g h^2}{2} - \frac{\beta g h^2 h_x^2}{4} + \{ h_t + [h \bar{u}]_x \} \tilde{\phi}$$

- New expressions for the kinetic and potential energies:

$$\frac{\mathcal{K}'}{\rho} = \frac{h \bar{u}^2}{2} + \frac{(2 + 3\beta) h^3 \bar{u}_x^2}{12}, \quad \frac{\mathcal{V}'}{\rho} = \frac{g h^2}{2} \left(1 + \frac{\beta h_x^2}{2} \right)$$

Derivation of a consistent mSerre model - II

Governing equations

- The Euler–Lagrange equations:

$$\delta\tilde{\phi}: 0 = h_t + [h\bar{u}]_x,$$

$$\delta\bar{u}: 0 = h\bar{u} + \tilde{\phi}h_x - \left[h\tilde{\phi} \right]_x - \left(\frac{1}{3} + \frac{1}{2}\beta \right) \left[h^3 \bar{u}_x \right]_x,$$

$$\delta h: 0 = \frac{1}{2}\bar{u}^2 - gh - \tilde{\phi}_t + \tilde{\phi}\bar{u}_x - [\bar{u}\tilde{\phi}]_x$$

$$+ \left(\frac{1}{2} + \frac{3}{4}\beta \right) h^2 \bar{u}_x^2 - \frac{1}{2}\beta gh h_x^2 + \frac{1}{2}\beta g[h^2 h_x]_x$$

- After eliminating the velocity potential:

$$\tilde{\phi}_x = \bar{u} - \left(\frac{1}{3} + \frac{1}{2}\beta \right) h^{-1} \left[h^3 \bar{u}_x \right]_x,$$

$$\tilde{\phi}_t = \frac{1}{2}\bar{u}^2 - \bar{u}\tilde{\phi}_x - gh + \left(\frac{1}{2} + \frac{3}{4}\beta \right) h^2 \bar{u}_x^2 + \frac{1}{2}\beta gh[h h_x]_x$$

- These equations have the full set of conservation laws:

$$\begin{aligned} \partial_t & \left[\frac{1}{2}h\bar{u}^2 + \frac{1}{12}(2+3\beta)h^3\bar{u}_x^2 + \frac{1}{2}gh^2 + \frac{1}{4}\beta gh^2 h_x^2 \right] \\ & + \partial_x \left[\left(\frac{1}{2}\bar{u}^2 + \frac{1}{12}(2+3\beta)h^2\bar{u}_x^2 + gh + \frac{1}{3}h\Gamma \right) h\bar{u} \right] = 0 \end{aligned}$$

Steady motion - I

Travelling wave solutions: periodic and solitary waves

- The mass conservation yields:

$$\bar{u} = -c d / h$$

- After substituting it in momentum conservation and some algebra:

$$\left(\frac{dh}{dx} \right)^2 = \frac{\mathcal{F} - (1 + \mathcal{C}_2 + 2\mathcal{F})(h/d) + (2 + 2\mathcal{C}_1 + \mathcal{F})(h/d)^2 - (h/d)^3}{\left(\frac{1}{3} + \frac{1}{2}\beta\right)\mathcal{F} - \frac{1}{2}\beta(h/d)^3}$$

Solvable analytically with **Jacobi elliptic functions!**

- For solitary waves $\mathcal{C}_1 = \mathcal{C}_2 = 0$:

$$\left(\frac{d\eta}{dx} \right)^2 = \frac{(\mathcal{F} - 1)(\eta/d)^2 - (\eta/d)^3}{\left(\frac{1}{3} + \frac{1}{2}\beta\right)\mathcal{F} - \frac{1}{2}\beta(1 + \eta/d)^3}$$

Solvable in **elementary functions** in parametric form!

Steady motion - II

Solitary wave solutions

- Introduce a new independent variable:

$$x(\xi) = \int_0^\xi \left| \frac{(\beta + 2/3)\mathcal{F} - \beta h^3(\xi')/d^3}{(\beta + 2/3)\mathcal{F} - \beta} \right|^{1/2} d\xi'$$

- The analytical solution:

$$\frac{\eta(\xi)}{d} = (\mathcal{F} - 1) \operatorname{sech}^2\left(\frac{\kappa \xi}{2}\right), \quad (\kappa d)^2 = \frac{6(\mathcal{F} - 1)}{(2 + 3\beta)\mathcal{F} - 3\beta}$$

- Wave-amplitude relation:

$$a \equiv \eta(0) = (\mathcal{F} - 1)d$$

How to choose the free parameter β ?

- The same *linear dispersion relation* analysis applies here!
- ... new ideas?

Nonlinear dispersion relation analysis

Some considerations towards the choice

- By following the classical work of McCowan (1891) [11]:

$$\frac{c^2}{gd} = \frac{\tan(\kappa d)}{\kappa d} = 1 + \frac{(\kappa d)^2}{3} + \frac{2(\kappa d)^4}{15} + \frac{17(\kappa d)^6}{315} + \frac{62(\kappa d)^8}{2835} - \dots$$

- The same result for the mSerre system:

$$\frac{c^2}{gd} = \frac{2 - \beta(\kappa d)^2}{2 - (\frac{2}{3} + \beta)(\kappa d)^2} \approx 1 + \frac{(\kappa d)^2}{3} + \left(\frac{1}{3} + \frac{\beta}{2}\right) \frac{(\kappa d)^4}{3} + \dots (*)$$

Meromorphic interpolation scheme:

- $\tan(z)$ is meromorphic with single poles at $\kappa d = \pm\pi/2, \pm 3\pi/2, \dots$
- $(*)$ has a single pole at $\kappa d = \pm\sqrt{6/(2+3\beta)}$
- The first poles at $\kappa d = \pm\pi/2$ coincide if

$$\beta = \frac{2}{3} \left(12\pi^{-2} - 1 \right) \approx 0.1439$$

Highest wave

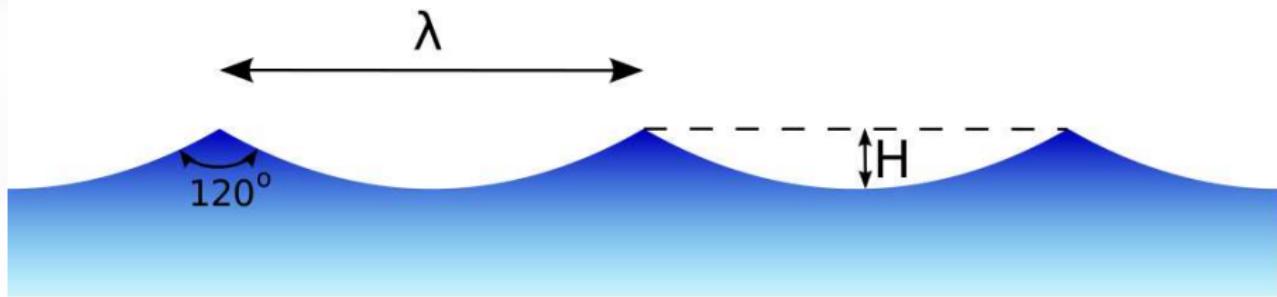
Sir George Gabriel Stokes (1819 – 1903)

In 1880 Stokes showed [12]:

- Existence of a limiting wave
- The angle at the crest is 120°

Reference (see Appendix B)):

Stokes, G.G. (1880). *Supplement to a paper on the theory of oscillatory waves*. Math. Phys. Pap., 1, 314–326.



Essentially unknown F. Serre's result

Determination of the highest solitary wave [13]

JUILLET-AOUT 1956 - N° 3

LA HOUILLE BLANCHE

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Contribution à l'étude des ondes longues irrotationnelles

Contribution to the study of long irrotational waves

PAR F. SERRE

INGÉNIEUR DOCTEUR
INGÉNIEUR A L'OMNIUM FRANÇAIS D'ÉTUDES ET DE RECHERCHES

Essentially unknown F. Serre's result

Determination of the highest solitary wave [13]

Le nombre de Froude F_0 , correspondant à l'onde solitaire limite, est donné par :

$$F_0^2 = \frac{\tanh 2 \sigma_0}{2 \sigma_0} ;$$

on a :

$$1,765558 < F_0^2 < 1,76707$$

d'où $F_0^2 = 1,766$ avec trois décimales exactes; la profondeur d'eau limite à la crête s'en déduit

par :

$$\frac{y_c}{y_0} = \zeta_c = 1 + \frac{F_0^2}{2} = \underline{\underline{1,883}}.$$

Remarks:

- Result obtained without any computers
- All digits are **correct!**

Almost-highest gravity waves on water of finite depth

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(Received 5 April 2000; revised 3 April 2001)

Almost-highest waves

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Table 8. *Properties of the highest solitary waves*

	Current results	Williams (1981)	Evans & Ford (1996)	Zhitnikov (2000)
\bar{c}_{sol}^*	1.2908904558	1.290889	1.29089053	1.290890455863
\bar{H}_{sol}^*	0.8331990844	0.833197	0.833199179	0.833199084520
\bar{M}_{sol}^*	1.970320658	1.970319	1.97032019	1.970320660132
\bar{I}_{sol}^*	2.543468132	2.543463	2.54346767	2.543468135155
\bar{C}_{sol}^*	1.714569239	1.714569	1.71456873	1.714569240534
\bar{T}_{sol}^*	0.535008835	0.535005	0.535008913	0.535008835971
\bar{V}_{sol}^*	0.437672693	0.437670	0.437672702	0.437672693444
$\bar{V}_{sol} + \bar{T}_{sol}^*$	0.972681528	0.972675	0.972681614	0.972681529415

- Taylor expansions around the crest yield

$$x = \left[\frac{(2+3\beta)\mathcal{F} - 3\beta\mathcal{F}^3}{(2+3\beta)\mathcal{F} - 3\beta_2} \right]^{\frac{1}{2}} \xi + \mathcal{O}(\xi^3), \quad \eta = a \left[1 - \frac{1}{4} (\kappa\xi)^2 \right] + \mathcal{O}(\xi^4)$$

- Peaked crests occur when $dx/d\xi = 0$ at $\xi = 0$. It happens if

$$\mathcal{F}^2 = 1 + \frac{2}{3\beta}$$

- The slopes at the crest are:

$$\left. \frac{d\eta}{dx} \right|_{x=0^\pm} = \left. \frac{d\eta}{d\xi} \right/ \left. \frac{dx}{d\xi} \right|_{\xi=0^\pm} = \mp \frac{\kappa a \sqrt{\mathcal{F}^2 + \mathcal{F} + 1}}{\sqrt{3}\mathcal{F}} = \mp \frac{\sqrt{2/3\beta} a}{a + d}$$

We can compare with the **full Euler prediction!**

Comparison of various choices of the free parameter β

Multiple choices are possible

Full Euler prediction:

$$\mathcal{F} \approx 1.290, \quad a/d \approx 0.8331, \quad \theta = 120^\circ$$

- Linear dispersion relation ($\beta = 2/15$, Taylor correct to $\mathcal{O}((kd)^4)$):

$$\mathcal{F} = \sqrt{6} \approx 2.45, \quad a/d = \sqrt{6} - 1 \approx 1.45, \quad \theta \approx 28.4^\circ$$

- Meromorphic interpolation scheme ($\beta = \frac{2}{3}(12\pi^{-2} - 1)$):

$$\mathcal{F} \approx 2.37, \quad a/d = \sqrt{6} - 1 \approx 1.37, \quad \theta \approx 37.3^\circ$$

- Inner angle interpolation scheme:

$$\beta = \frac{4\beta + \sqrt{2\beta} - 1}{3\beta - \sqrt{8\beta} + 2} \implies \beta \approx 0.3456$$

$$\mathcal{F} \approx 1.711, \quad a/d \approx 0.711, \quad \theta \approx 120^\circ$$

< 15% relative error in amplitude!

Instead of conclusions

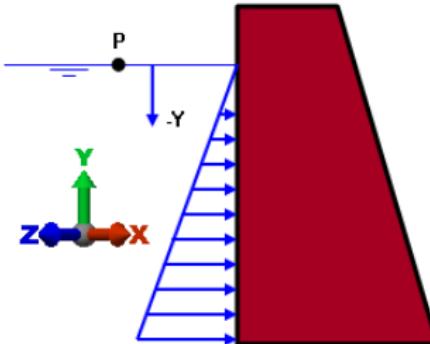
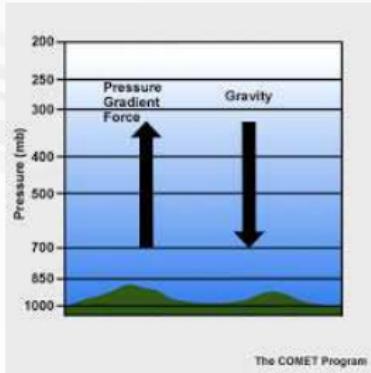
A note of caution

Attention, these are not synonyms:

Dispersive \neq Non-hydrostatic

Counter-examples can be constructed [15]:

There are **non-hydrostatic** models, which are **non-dispersive!**



Thank you for your attention!



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